

# GEOLOGIC COLUMN AND UNIT DESCRIPTIONS

AGE	ROCK UNIT	LITHOLOGY; THICKNESS WHERE KNOWN	REMARKS	ECONOMIC VALUE	REFERENCES	
QUATERNARY	Alluvium	Sand, gravel, and clay. Thickness less than 10 meters.	Recent deposits of sand, gravel and clay covering river flats.	Placer gold was worked in ancient times from the river gravel at various places.	ASANO, Gorō, 1939, Geology of the district of Ja-shui-tang (夹山水) hot spring, Chien-chang Hsien (乾岔县), Jehol: Bull. Geol. Inst. South Manchuria Railway Company, no. 95.	
	Diluvium	Aeolian sand, loess, and gravel. Thickness less than 40 m.	Pleistocene deposits including aeolian sand, loess, redeposited loess, and gravel.	Placer gold was worked until quite recently, especially from the terrace deposits along the streams of the gneiss region.	AZUKI, Toshiyoshi, 1936, Origin of the nitro deposits in the vicinity of Hsueh-shui (薛水), Hsü-chui (许水), Jehol: Bull. Geol. Inst. South Manchuria Railway Co., no. 87.	
TERTIARY	Unconformity					
	Neogene basalt	Augite-olivine basalt, tuff, and breccia. Thickness 100 m to 400 m.	Neogene basalt consists mainly of flows and sheets of augite-olivine basalt intercalated with tuff and breccia. It is best developed in the mountain land around Hsi Shan (西山), west of Ling-yuan (凌源). Age of its eruption is believed to be Pliocene.		BESSHO, Bunichi, 1936, Geologic report of coal fields around the district of Chih-feng (赤峰): Unpublished report furnished by Geol. Inst. South Manchuria Railway Co. for official use only.	
	Effusive contact					
	Rhyolite	Rhyolite, trachyandesite, tuff, and breccia. Thickness more than 300 m.	Rhyolite of probably Middle Cretaceous age is mostly in flows and dikes, associated with tuff and breccia. Trachyandesite also occurs though in minor amount. Because of scarce data, the Cretaceous rhyolite is hardly distinguishable from the rhyolitic volcanic series of the Upper Jehol formation.		1931, Report on the nitro deposits in the vicinity of Hsi-shui, Jehol: Unpublished report furnished by Geol. Inst. South Manchuria Co. for official use only.	
	Effusive contact					
MESOZOIC	Quartz porphyry	Quartz porphyry and granite porphyry.	Quartz porphyry and granite porphyry intruded during the period of Lower Cretaceous; may have been differentiated from the same magma as that of the granite (g <sub>2</sub> ).	(gold mines) 1. Chang-kou (昌古); 2. Hsü-tien (许田); 3. Tai-ho-kou (台河口); 4. Wei-lin-kuo (魏林郭); north of Fu-chia-wo-pai (富家窝排); 5. Tsao-mao-shan (赵毛山); 6. Kao-shia-tien (高家田); 7. A-ho-yung-tau (阿合营套); 8. Chih-chang (赤峰); 9. Hsiang-chang-yung-tau (香厂营套); 10. Ching-chien-tai (荆钱台); 11. Wu-lung-tai (五龙台); northwest of Ching-chien-tai; 12. Hsiao-chang-kou (小厂口); north of Fu-lung-shan (伏龙泉); 13. Tso-shang (左乡); 14. Chien-chang-yung-tau (锦厂营套); south of Lung-wang-miao (梁王庙).		
	Cretaceous granite	Biotite granite, quartz monzonite, and felsitic granite.	Biotite granite, quartz monzonite and felsitic granite. The age of intrusion of these rocks may be late Jurassic or early Cretaceous, from the fact that in the mining localities (see Economic Value) a number of ore veins which are genetically related with the granitic intrusives cut the flows and sheets of andesite and rhyolite of the Jehol formation.	(gold-copper mine) 15. Ping-chuan (平川)		
	Diorite	Diorite and granodiorite.	Diorite and granodiorite; may be products of differentiation of the same magma as that of the granite (g <sub>2</sub> ).	(silver-lead mines) 1. Yang-shu-lin (杨树林); 2. Yu-shu-ti (榆树台); 3. San-tao-kou (三套口)		
	Intrusive contact					
	Jehol formation	Upper: alternation of tuff, sandstone, and conglomerate, with acidic volcanic rocks and pyroclastic rocks. Thickness about 500 m.  Middle: sandstone, shale, conglomerate, and coal. Thickness 500 m to 1,000 m.	The Jehol formation is predominantly Jurassic accompanied locally by Uppermost Triassic and Lower Cretaceous; corresponds to the so-called "Jehol formation" by Iino MATSUZAKI and others, and is divided into the upper, the middle, and the lower parts as follows:  The Upper Jehol formation consists of an alternation of tuff, sandstone and conglomerate, and is characteristically intercalated with flows and sheets of acidic volcanic rocks and pyroclastic rocks. Its total thickness in the district of Ling-yuan may be more than 500 m. With no data available it is often very hard to distinguish the Upper Jehol formation from the overlying rhyolite flows; this is particularly so in the northeastern part of the map area. The age of the Upper Jehol formation is considered to be Uppermost Jurassic, although locally it may range to Lower Cretaceous.  The Middle Jehol formation consists of sandstone, shale, and conglomerate, and is characterized by the occurrence of fresh-water fossils such as <i>Esteria</i> and <i>Lyocoptera</i> . The Middle Jehol formation, distributed in the districts of Ling-yuan, Ku-tai-hai (库台海), Ta-cheng-tau (塔城套), and in the area east of Wang-shu-kou (王书口), is characterized by oil shale-bearing beds, while the exposures in the vicinities of Su-tai-kou (苏台口), Wu-chia (五家), Wei-tan-kuo (魏田郭), Ta-tai-hsi (塔台西), Pa-li-chien-tien-tau (八里店套), Su-lung-kou (苏龙口) southeast of Wu-fang-tien (五方甸), and in the area southeast of Chien-chang (锦昌) are marked with a coal-bearing formation. The Middle Jehol formation generally lacks tuff and contemporaneous volcanic rocks which present a characteristic feature of the Upper Jehol formation. A thick basal conglomerate occurs. The total thickness may range between 500 m and 1,000 m. Fossils reported from the <i>Lyocoptera</i> -bearing formation near Iao-shang-tau (腰厂套) southeast of Ling-yuan were: <i>Non-parasurus splendens</i> Endo, <i>Hymenoceras orientalis</i> Endo et Shikama, <i>Lyocoptera davidi</i> (Sawagoe), <i>Sinobolinitis latyungensis</i> Ping, <i>Ephemeropsis bristalli</i> Rickward, <i>Astracis ticinii</i> V. Strahlen, <i>A. spinirostris</i> Luksun, <i>Coskanowakia rigida</i> Heer, and <i>Schizophoria koslowi</i> Tabe et Endo. Those from Ta-shan-fang-tau (塔山房套) north of Ling-yuan were <i>Lyocoptera jeholensis</i> Grabau, <i>L. jeholensis</i> var. minor Grabau, <i>Esteria</i> sp., and <i>Corbicula jeholensis</i> Grabau. Ta-cheng-tau (塔城套) is also a well-known locality of the <i>Lyocoptera</i> -bearing formation. Fossils reported from the coal-bearing formation of Wu-chia coal field were <i>Platoceras</i> sp., <i>Phoenicopsis</i> sp., <i>Gingkotites sibirica</i> Heer, <i>Coskanowakia dichotoma</i> Heer, <i>Pityophyllum</i> sp., <i>Samarura</i> sp., and a bone of <i>Lyocoptera</i> (?). As to the age of the Middle Jehol formation no decisive agreement has been reached among geologists, and it varies between Middle Jurassic and Uppermost Jurassic.	(coal) 1. Wu-chia (五家); 2. Su-tai-kou (苏台口), southwest of Tung-feng-tai ( tung 风台); 3. Wei-tan-kuo (魏田郭), northwest of Wei-tan-kuo (魏田郭); 4. Yang-chia-kou (杨家口); 5. Lao-yao-kou (老姚口), northeast of Wu-fang (五方); 6. Ta-yu-hsi (塔峪)		
Unconformity						
PALEOZOIC	Lower andesite, andesitic pyroclastic rocks, sandstone, conglomerate, and coal.	Lower: andesite, andesitic pyroclastic rocks, sandstone, conglomerate, and coal. Thickness more than 500 m.	The Lower Jehol formation is generally characterized by a predominance of andesite and andesitic pyroclastic rocks, and is covered either by the Middle Jehol formation or by the Upper Jehol formation. The succession of the Lower Jehol formation at the Ni-yung-tau (牛营套) coal field, 20 km south of Ling-yuan, is reported to be as follows, in descending order: a reddish formation (consists of conglomerate, sandstone, andesitic lava, agglomerate, and tuff), a coal-bearing formation, and a bedded, thick basal conglomerate. The Sianian limestone was thrust over the reddish formation from northwest to southeast, and the limestone is in turn unconformably covered by the acidic volcanic series of presumably Upper Jehol formation. Fossils such as <i>Gladophlebia halburgenis</i> , and <i>Neocalamites</i> sp. were reported from the coal-bearing formation. The age of the Lower Jehol formation varies between Lower Jurassic to Rhaetic according to different opinions among geologists.	(coal) 7. Ni-yung-tau (牛营套)  The Ni-yung-tau coal field was intermittently operated by the neighboring villagers for local consumption. The mine locality is hard to reach. The coal-bearing beds belong to the Lower Jehol formation.		
	Andesite	Andesite, and agglomeratic andesite.	Flows and sheets of andesite and agglomeratic andesite; mainly contemporaneous with the Lower Jehol formation, and partly associated with the Upper Jehol formation.			
	Unconformity					
PALEOZOIC	Permian formation	Slate, sandstone, conglomerate, and coal. Thickness about 450 m.	The Upper Paleozoic in the map area is represented by a Permian terrestrial formation consisting of slate, sandstone, conglomerate, and coal. The formation, along with the Lower Paleozoic formation, crops out in the districts of San-tao-kou (三套口), Ping-chuan Hsien, as a consequence of local erosion cutting through the overlying thrust. The total thickness of the formation, judged from that in the Lin-yu sheet (NK 50-12) area, may not be less than 450 m.			
	Unconformity					
PALEOZOIC	Cambro-Ordovician formation	Limestone, magnesian limestone, shale, and slate. Thickness more than 500 m.	The Lower Paleozoic formation consists of limestone, magnesian limestone, shale and slate. No fossil has been collected from this formation in the map area, but from the lithological resemblance it can be considered as a northern extension of the Cambro-Ordovician formation of the Lin-yu sheet area. The total thickness may be more than 500 m.			
	Unconformity					
PRECAMBRIAN	Sinian system	Upper: limestone, magnesian limestone, siliceous limestone, siliceous phyllite, calcareous slate, and red slate. Thickness about 2,000 m.  Lower: quartzite, sandstone, and conglomerate. Thickness 1,000 m to 1,500 m.	The Upper Precambrian in the map area is represented by the Sinian system. Stratigraphy of this system in the map area has not been studied in detail, but generally it shows no large difference from that of the adjacent Lin-yu sheet area. Lithologically the Sinian system can be roughly divided into the upper and lower formations.  The upper formation with a probable thickness of more than 2,000 m has the following succession, in descending order: a limestone, a magnesian limestone, a siliceous limestone, an alternation of siliceous phyllite and calcareous slate, and a red slate at the base.  The lower formation consists, in descending order, of a thickly bedded quartzite, a siliceous sandstone, a conglomerate, and a basal conglomerate. Thickness ranges between 1,000 m and 1,500 m.  No recognizable unconformity occurs between the upper and lower formations. Generally a few beds of red slate mark the base of the upper formation. The Sinian system in this map area presents large-scale varying structures associated with thrusts.			
	Unconformity					
	Wufai system (?)	Schist complex consisting of crystalline limestone, mica schists, and hornfels. Thickness unknown.	Schist complex of the Wufai system (?); consists of crystalline limestone, mica schist, and hornfels. It crops out in the districts of Yu-shu-ti (榆树台), Chih-chiang (赤峰), Hsiang-yeh-tai (香厂台), and Ching-kuo (庆郭). It is a product of contact metamorphism caused by injection of granitic rock or orthogneiss into sedimentary rocks. The age of the schist complex has not been determined yet, but it is generally correlated with the Wufai system or Lower Precambrian.			
Unconformity						
Sangkan system (?)	Gneiss complex consisting of orthogneiss, injection gneiss, and schist. Thickness unknown.	Gneiss complex of the Sangkan system (?); consists mainly of medium-grained orthogneiss and injection gneiss of variable mineral compositions. Some crystalline schist also occurs. It is a complicated mixture formed by metamorphic diffusion of both the pre-existing sedimentary rocks and the later granitic intrusives. The age of this complex may be Precambrian. From its lithological resemblance, G. B. Barbour correlated the gneiss complex of this map area with the Sangkan system of Kaigan area.				

(Column not drawn to scale)